

Flood transported quartzites: Part 3— failure of uniformitarian interpretations

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The distribution of quartzite gravel, cobbles and boulders on the mountaintops, ridges plateaus and valleys of northwestern USA and southwest Canada presents a serious challenge to the uniformitarian paradigm. Such geologists have recognized the inadequacy of present day processes to explain these deposits. Of all the uniformitarian hypotheses proposed, the only one that does not seem to have been eliminated is the ‘outrageous’ hypothesis of large-scale fluvial transport. This hypothesis still falls far short of accounting for the geological evidence, including the extensive distribution of the quartzite deposits, the long distance of transport, the iron staining of the quartzite clasts and their abundant percussion marks.

We have described many of the quartzite gravel locations that we have observed or noted in the literature in previous papers.^{1,2} The likely sources of quartzite are in the Rocky Mountains of northwestern Montana, central and northern Idaho and adjacent Canada, mainly in the Belt Supergroup. (Uniformitarian names are used for communication purposes only and are not meant to imply acceptance of the claimed ages or an absolute, but compressed, geological column and timescale.)

Paleocurrent directions and a general decrease in clast size demonstrate that the quartzites were eroded from the northern Rocky Mountains. The quartzite material becomes well rounded the farther away they are from their source areas. Gravel deposits are found on mountaintops, ridges, plateaus and in valleys at lower elevations. They have been transported over 1,000 km to the east and 700 km to the west over what appears to have been generally flat ground.

The quartzite gravels range in thickness from a thin lag to deposits up to 4,750 m in northeast Idaho and over 3,000 m east and northeast of Jackson Hole, Wyoming. There is a general size decrease in the gravel away from the sources. The quartzite gravel is often iron stained and commonly exhibits percussion marks. Thick deposits, occurring mainly in deep, likely fault-controlled valleys or basins, often have pressure solution marks and are fractured. Gold is sometimes found within the matrix of the gravels, for instance in central and northeast Oregon and northwest Wyoming. The observation and analysis of these quartzite gravels represents the culmination of a fifteen-year project. Yet, much remains to be learned from these and similar deposits.

Uniformitarian explanations for far-traveled quartzites

How do uniformitarian geologists explain all of these quartzite gravels that were transported so far from their sources? Uniformitarian scientists have observed these quartzite gravels for many years, although they have been reluctant to publish much information on them. They have several standard explanations when attempting to account for the quartzite gravel. We show that these uniformitarian

explanations are not supported by the data. We start with the plains east of the continental divide

Plains east of continental divide

On the plains east of the continental divide, the distance from source rocks and the generally flat planation surface remnants upon which the quartzite gravels lie are specific issues that need to be explained. However, we will not address the question of the formation of planation surfaces since it has been discussed elsewhere.²⁻⁸

Uniformitarian scientists have attributed the widespread quartzite gravels spread far to the east to river transport confined to valleys.^{9,10} Since streams would form the lower parts of the landscape and the gravels are now found on plateaus, the gravels suggest *relief inversion*, in which erosion was faster along the interfluvial ridges over millions of years. However, rivers cannot transport gravels many hundreds of miles from their source over such gentle slopes.³

The inadequacy of river transport was also recognized by Leckie and Cheel.¹¹ Instead, they proposed that multiple braided streams spread a large sheet of gravel 150 km out from the Rocky Mountains in a semi-arid climate, forming what is called a ‘braidplain’.¹² Figure 1 shows their interpretation of this braidplain, which extended so far east as to cover the incipient Sweetgrass Hills and Bears Paw Mountains. Then the tectonic uplift of the Sweetgrass Hills and Bears Paw Mountains provided an ‘eastward boost’ to the gravel (figure 2). The boost was partially provided by rivers or streams in confining valleys. Subsequent erosion over about 30 million years produced isolated gravel-capped erosional remnants at four general levels.

There are a number of problems with Leckie and Cheel’s hypothesis.³ Any mechanism for deposition of the Cypress Hills Formation, Wood Mountain Plateau gravel, and Flaxville gravel must be able to account for:

- The laterally extensive, planar, surficial planation surfaces at generally four elevations,
- The laterally extensive sheets of coarse gravel that mantle the planation surfaces,
- The transport of exotic quartzites 300 to 700 km from their nearest source over slopes of less than 0.1 degree.

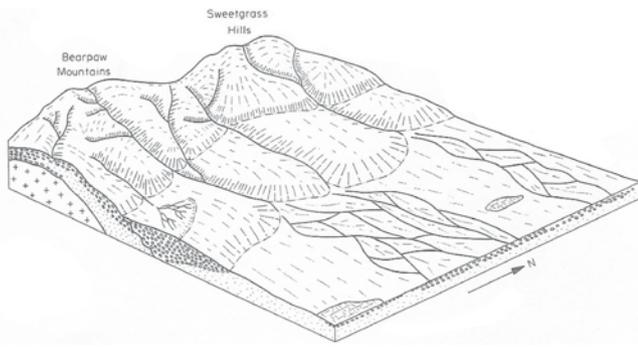


Figure 1. Proposed braidplain formed in a semi-arid environment up to 150 km east of the Rocky Mountains. Note that this braidplain spread over the incipient Sweetgrass Hills and Bears Paw Mountains of north central Montana. (From Leckie and Cheel,¹¹ p. 1,928).

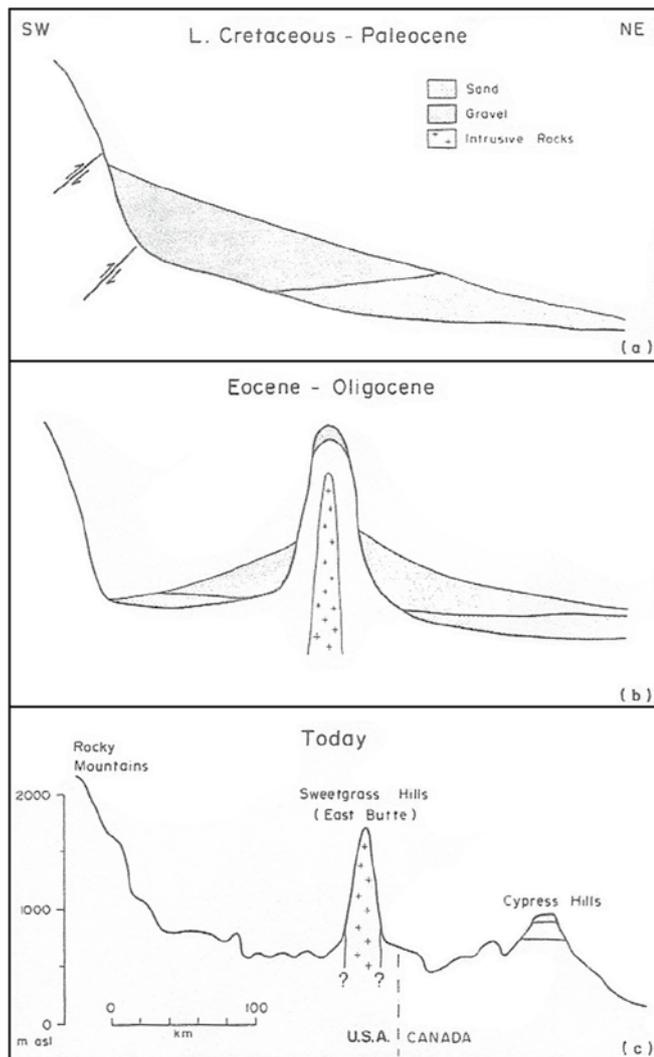


Figure 2. Sequence of events proposed for the origin of the Cypress Hills quartzite gravels. (a) Coarse quartzite debris shed eastward off the Rocky Mountains. (b) Granitic intrusive rocks uplifted previously deposited sediments. (c) Uplifted sediments recycled northeastward to their present location in the Cypress Hills followed by erosion of the whole area. (From Leckie and Cheel,¹¹ p. 1,929).

Furthermore, Leckie and Cheel¹² even suggest that the quartzites were transported from central Idaho, a distance 200 km farther to the southwest than the northern Rocky Mountains of western Montana. If we include the occurrences of quartzite gravel in North Dakota, the transport distance is over 1,000 km. These distances seem much too far for such an elaborate mechanism as proposed by Leckie and Cheel.¹²

The quartzite gravels show no evidence of being deposited in channels or as alluvium. Transport by 'palaeorivers' would have eroded and incorporated a mixture (probably a majority) of local lithologies, but we observe that 90% or more of the gravels are quartzite.

Klevberg considered many possible transport mechanisms from the Rocky Mountain Front far onto the high plains.¹³ Various types of mass flows, including hyperconcentrated flows, were considered in his analysis. There are many problems in understanding and much debate over the nature of mass movement.¹⁴ One of the main problems is that scientists attempt to understand the flow process just from the deposit.¹⁵ Furthermore, the classification system is confused and changes frequently. For example, there is much confusion over the definition of a turbidity current and the products of a turbidity current.¹⁶⁻¹⁸ True turbidity currents are shown to contain less than 9% sediments. This is probably why true turbidity currents are not very erosive. Hyperconcentrated flows are generally similar to debris flows. However, there is also confusion over the meaning of a hyperconcentrated flow, which seems to contain less sediment than most debris flows.¹⁹

Based on the observations, the evidence overwhelmingly points to transport by a sheet flow of water. Flash floods and the break up of ice dams during the Ice Age can be eliminated.³ Furthermore, the Cypress Hills Formation cobbles and boulders were at first spread over an area much more extensive than the Cypress Hills.⁴ After deposition, the high plains continued to be eroded. Generally three more planation surfaces formed on the high plains with a total depth of erosion of 760 m down to the level of the rivers surrounding the Cypress Hills. This is the signature of sheet erosion by water that becomes more channelized with time.

Klevberg used the Chezy Equation for coarse sediment bedload transport and the Keulegan Equation for bed roughness.¹³ From the fact that percussion marks are abundant in the gravels, Klevberg reasoned that the quartzites were in suspension briefly in fast turbulent flow. There is a relationship between the horizontal and fall velocity of suspended clasts that Klevberg used to calculate current velocities near the bottom of the flow. The maximum size and shape of the clast briefly suspended in turbulent flow was considered. A 15-cm wide bullet shaped clast was considered the maximum. The spin of the non-spherical clast was also considered in the calculation. The results showed that transport of the gravel in brief suspension would require transport by a current with a depth of at least 55 m

Table 1. Minimum flow velocity and depth for two thresholds and shapes of quartzite clasts. (After Klevberg and Oard³).

Estimates of current speed and depth for suspended rocks		
Rock size and shape	Minimum current speed	Minimum current depth
10 cm sphere	30 km/h	15 m
15 cm wide bullet	105 km/h	55 m

and moving with a velocity of at least 108 km/h (figure 3 and table 1).³ While the data upon which these calculations are based provide considerable spread, resulting from the many differences in bed roughness, clast shape and other variables, it should be noted that the calculations provide *minima* for current parameters. The inferred velocity is several times faster than is typical of flash floods on earth that travel down steep mountain slopes. However, the spread of Cypflax was over a low slope and a large area.

The Sweetgrass Hills (figures 4 and 5) are 100 to 150 km from the Rocky Mountain front and are small crystalline rock intrusions, called stocks.²⁰ It is hard to imagine that these isolated hills that range up to about 2,100 m ASL (above sea level) and up to 1,000 m above the plains could shift huge amounts of quartzite another 150 km northeast to form a sheet deposit of gravel on the Cypress Hills Plateau. How does the Leckie and Cheel hypothesis account for the Swift Current Creek, Wood Mountain Plateau and Flaxville Plateaus, a few hundred km farther east that are also capped by quartzite gravels? What about the quartzite gravels in North Dakota? There is very little quartzite gravel around the Sweetgrass Hills, although we found an isolated cobble above the depth of glaciation while climbing West Butte.

The Bears Paw Mountains are about 200 km from the Rocky Mountain front to the east, farther than Leckie and Cheel's¹² 150 km of transport in their initial braidplain. (It is possible that quartzite gravel could have originated from quartzite outcrops in the Little Belt and Big Belt Mountains to the south and south-southwest, which are closer than the Rocky Mountains, but the volume of quartzite in these mountains is limited and paleocurrent directions in the quartzite gravels indicate more of a westerly source). Furthermore, the quartzites that outcrop in the Rocky Mountains but east of the divide, especially south of Glacier National Park, are of lower grade (less metamorphosed) than the Cypflax gravel. Therefore, the Cypflax gravel likely originated west of the continental divide! This lower grade quartzite is generally what forms the non-Cypflax gravels south of the Cypflax exposures.

Considering the many problems with the braidplain/tectonic boost hypothesis, it seems like a very speculative and *ad hoc* concept. Leckie and Cheel even admit that sedimentologically the Cypress Hills Formation is poorly understood.²¹

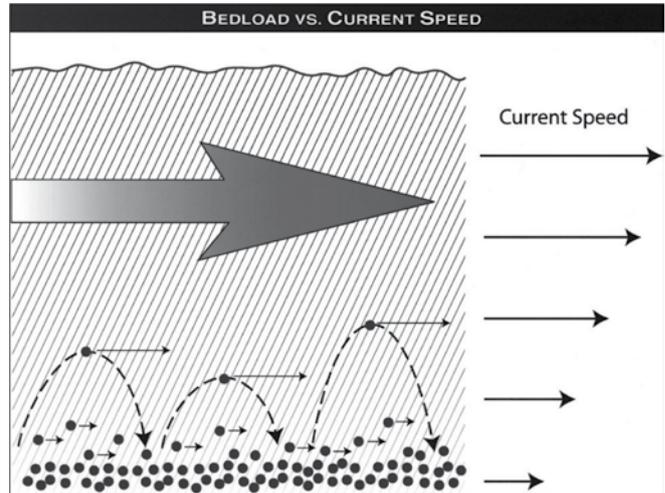


Figure 3. Schematic showing how we believe the percussion marks formed in the Cypress Hills gravel. A current will carry bedload dragged along the bottom, smoothing the surface into a planation surface. If currents are fast enough, clasts will be briefly carried up into suspension within the current. As the clasts crashed down into the bedload, percussion marks were formed.



Figure 4. West Butte of Sweetgrass Hills, north-central Montana.



Figure 5. Haystack Butte (left) and Middle Butte (right) of the Sweetgrass Hills, north-central Montana.

Southwest Montana, northwest Wyoming and adjacent Idaho

There are several hypotheses to account for the transport of gravel into northwest Wyoming.²² The four main hypotheses are:

1. Derivation from local quartzite sources
2. Transport by tectonic uplift in eastern Idaho^{21,23}
3. Transport on thrust sheets that have moved east for a long distance^{24,25}
4. Spreading by alluvial fans and/or rivers with progressive eastward recycling²⁶⁻²⁹

Local sources can easily be eliminated because the amount of quartzite in local mountains is far too small, and the paleocurrent directions in the quartzite gravel indicate currents from the west.³⁰ Although there is a small amount of quartzite that outcrops in the Beaverhead Range, this source is excluded because the light-tan quartzite in that range is different from the multi-coloured quartzite clasts of the Beaverhead Formation.³¹ Furthermore, there are only three sandstone formations in the area that could have a little quartzite, but these sandstones do not resemble the quartzite gravels described in Part 1 of this article.¹ Love, a uniformitarian expert on Wyoming geology, concluded:

‘Immediately eliminated are the Paleozoic and Mesozoic rocks in mountains adjacent to Jackson Hole. The Flathead (Cambrian) and Tensleep (Pennsylvanian) Sandstones and the Quadrant (Pennsylvanian) Formation are the only ones in the Paleozoic sequence that are even moderately quartzitic, and this only in localized areas. These sandstones do not resemble, in physical

appearance, variety of colour or degree of metamorphism, the quartzites in the conglomerates. No widespread quartzitic sandstones of consequence are in the Mesozoic sequence.

‘There is no 500-sq-mi [1,300 km²] or larger area of Precambrian quartzite in the cores of the Wind River, Gros Ventre, Teton, Centennial, Madison, Gallatin, Beartooth or Washakie uplifts ... In fact, quartzite is an uncommon lithology.’³²

The second possibility for the source of the quartzites is the tectonic uplift of a block of Belt quartzite in eastern Idaho, called the Targhee uplift (figure 6), which was championed by Love.²³ Love did not believe rivers could transport quartzites very far, so he reasoned that the source had to be close to the west. The Targhee uplift is supposed to be the source of the Belt Supergroup quartzites that spread both northward to form the thick Divide quartzites and eastward by high-velocity rivers into the Jackson, Wyoming, area. Then the uplift subsided and was covered by the Snake River basalts. Lindsey concludes that there is *no evidence for the Targhee uplift*, and it has not been detected by geophysical studies underneath the eastern Snake River basalt plain.²² Moreover, the Targhee uplift cannot account for the more than 4,000-m depth of the Divide quartzite, mainly because the paleocurrent directions in these quartzite gravels are from the west or southwest, not from the south as would be required by the hypothesis. Janecke *et al.* also dismiss the Targhee uplift hypothesis.³³

The third hypothesis of transport on eastward moving thrust sheets does not fare any better. Janecke *et al.* write that quartzites in what they believe are thrust sheets in southwest Wyoming and adjacent southeast Idaho were covered by sediments at the time the quartzite gravel was deposited.³³ Furthermore, they state that these sources of quartzite for the Jackson Hole area were too small and too far away. Lindsey is likewise skeptical of the overthrust mechanism.³⁰

The elimination of the above mechanisms leaves only the fourth mechanism: transport by rivers and/or alluvial fans from the west. This mechanism seems to be the only hypothesis currently entertained by uniformitarian scientists.²⁷ Before we discuss this idea, we must know the source and distance of transport of the quartzite gravels and conglomerates. It has been admitted by all investigators that the nearest source of the quartzites is in central Idaho along the eastern flank of the Idaho Batholith (figure 7) about 200 to 350 km from the Jackson, Wyoming area.³⁴⁻³⁶ The quartzite gravels are similar to the Belt Supergroup quartzites that outcrop in central and

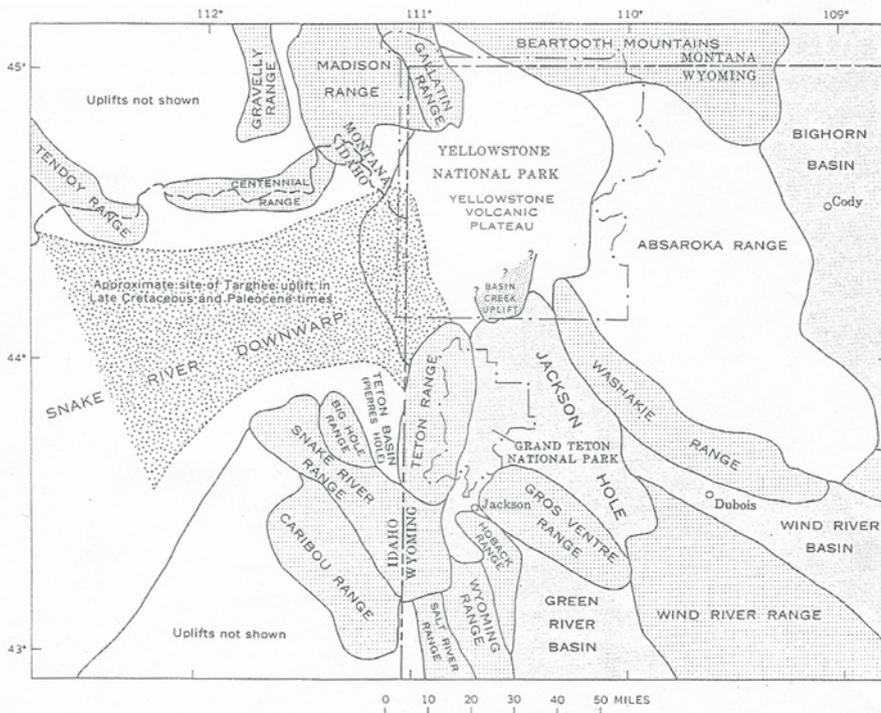


Figure 6. Postulated area of the Targhee uplift, which coincides with the Snake River basalts laid down in a downwarp. (From Love,²³ p. 4).

northern Idaho.³⁷ We think it likely that the quartzites could have come from the unroofing of the Idaho Batholith and the erosion of the Belt Supergroup in the area. Lindsey thought that the farthest transport could have been up to 450 km away to the west or northwest.³⁸ For the quartzites farther east in the western Bighorn Basin, Kraus believes these gravels were deposited in a ‘braidplain’, eroded from the quartzites in the Jackson area, 100 km to the west.^{28,29} But the quartzites also extend into the eastern Bighorn Basin, another 50 km east. Therefore, the total distance of transport ranges from 350 to 600 km!

Lindsey attempts to make a case from simple calculations that very large quartzite boulders can be transported up to 450 km away by rivers.³⁹ He also states that there is a lack of evidence for reworking, which would eliminate tectonic transport by the uplifting of mountain ranges, as envisioned by Leckie and Cheel for the Cypress Hills gravels discussed above.¹²

Recently, Janecke *et al.* have postulated that gravel

transport from west of Salmon, Idaho, was by two large ‘Eocene’ paleorivers.²⁷ These two paleorivers merged in southwest Montana and spread quartzites east and southeastward, forming what they call the Beaverhead-Harebell-Pinyon megafan that stretches from southwest Montana and adjacent Idaho into northwest Wyoming (figure 8). The paleovalleys are mainly *defined by the outcrops of quartzite*, some of which are on or near mountain ridges. Figure 9 shows a large distinctive quartzite boulder just west of Lemhi Pass, which is along the continental divide and is one of Janecke *et al.*’s paleovalley locations.²⁷ Uniformitarian scientists believe these ridge top gravels were tectonically uplifted after deposition in the paleoriver. These paleorivers are supposed to have transported several thousand cubic km of quartzite into northwest Wyoming from 200 to 350 km away!

We examined a number of outcrop locations mentioned by Janecke *et al.*²⁷ and found quartzite boulders up to almost 1 m in length. We also noted fewer rounded quartzites to the west, as would be expected if the source is generally west of Salmon, Idaho. We observed well-rounded quartzite gravel in several locations south and west of the postulated paleovalleys near and south of the Morgan Creek/Panther Creek divide, around 50 km southwest of Salmon, Idaho. Some of these quartzite gravels were reworked by local mountain glaciers during the Ice Age. One outcrop of boulders was in the headwaters of Morgan Creek about 30 km north of Challis, Idaho. The largest boulder at this locale was about 0.7 m in diameter, and some boulders had percussion marks (figure 10), a sign of very fast, turbulent flow. Therefore, the idea of two paleovalleys funnelling quartzites eastward is a simplification.

As mentioned in Part 1, the limestone conglomerate and quartzite gravels are usually separated. Ryder laments the fact that the quartzites rarely picked up limestone when traversing over mostly limestone terrain:

‘It still remains enigmatic how a monolithologic quartzite conglomerate could be transported across a predominately limestone terrain without becoming partially contaminated.’⁴⁰

The limestone conglomerate probably represents catastrophic local erosion as the mountains of southwest Montana and east-central Idaho rose upward. This is deduced from the abundance of *angular* clasts. Water was also involved since some of the clasts are rounded. One would certainly expect that there would be plenty of limestone clasts and limestone conglomerate mixed with the quartzites if these quartzite gravels were spread east by rivers. Could it be that the currents that eroded the quartzite from near the Idaho batholith were so fast that the limestone was rapidly pulverized?

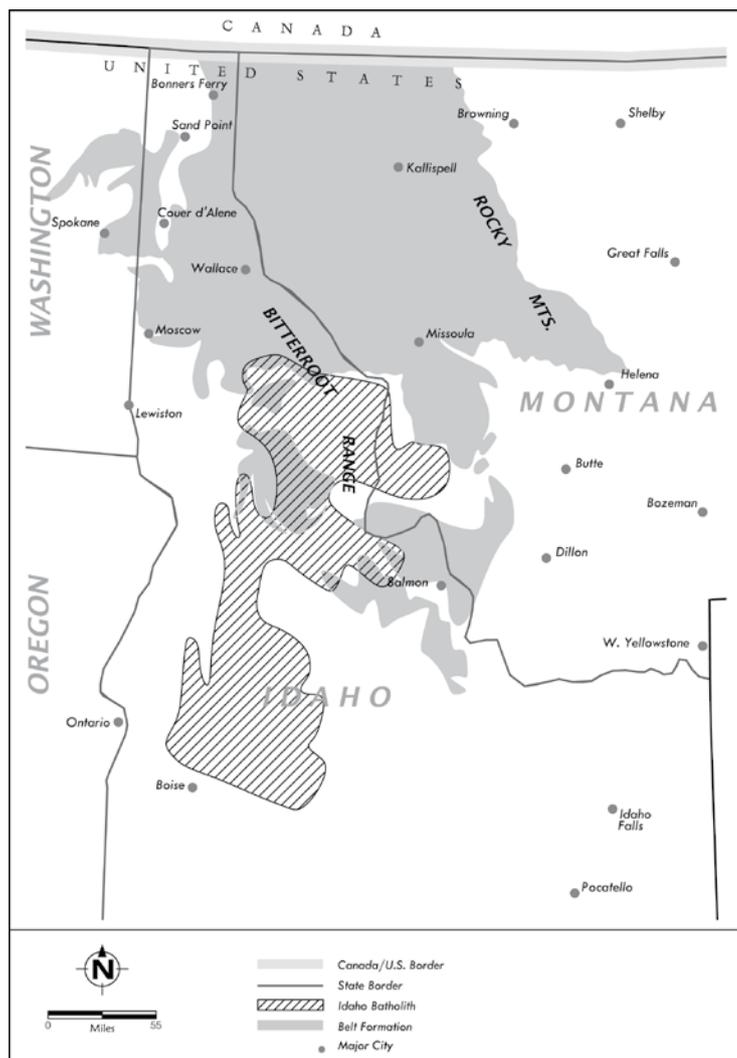


Figure 7. Outcrop areas for quartzite in northern and central Idaho and northwest Montana. Idaho batholith in west-central, Idaho, also indicated.¹

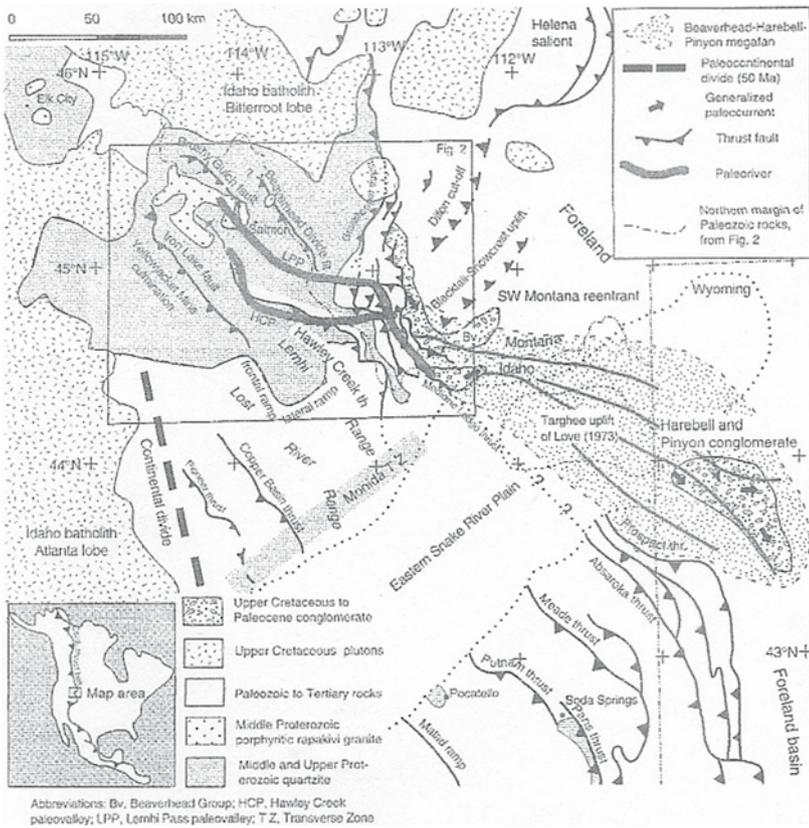


Figure 8. Postulated paleovalleys from near Salmon, Idaho, into southwest Montana forming the 'Beaverhead-Harebell-Pinyon megafan' that spread thick quartzite gravels into northwest Wyoming. (From Janecke *et al.*,²⁷ p. 439).



Figure 9. Large blue-banded quartzite from just west of Lemhi Pass, which is on the continental divide. This location is one of the paleovalley locations of Janecke *et al.*²⁷ Note that the boulder is subrounded, since it is closer to the source.

Can rivers really transport large quartzites from central Idaho to east of Jackson Hole, not to speak of the Bighorn Basin farther east? Janecke *et al.* do not provide any quantitative evaluation of their paleovalley hypothesis.²⁷ Their paper simply tries to trace out, based on the location of the quartzite gravels, the supposed routes of transport. Furthermore, we would expect that 'paleorivers' would have eroded and incorporated the local lithologies, but the clasts are more than 90% quartzite.

To provide the gravitational potential energy for



Figure 10. Large quartzite with percussion marks from the headwaters of Morgan Creek, about 30 km north of Challis, Idaho, and south of the paleovalley locations of *et al.*²⁷

transport by ancient rivers for quartzite gravels up to 25 cm in diameter, Lindsey calculated the height of the mountains in central Idaho as ranging up to 4,500 m above the level of the quartzites in northwest Wyoming with a 1° slope, his minimum slope estimate.²⁶ Since the Jackson, Wyoming area is around 1,800 m ASL, the mountains in central Idaho would have to have been 6,300 m ASL. This scenario is of doubtful validity, even for the lowest postulated slope by Lindsey.²⁶ There are many large boulders greater than 25 cm in diameter and there is little elevation difference between the source of the quartzites and the current locations of the quartzite gravels. The gravels in northwest Wyoming probably are of higher elevation. There would have to have been tectonic subsidence of the northern and central Rockies following the paleovalley river transport for the mechanism to be viable.

Washington and Oregon

The only hypothesis that has been entertained for the spread of quartzite gravels west of their Rocky Mountain source is transport by rivers.^{41,42} Researchers correctly deduced that the quartzites originated from the Belt Supergroup. The rivers are referred to as the 'ancestral' Columbia, Salmon and Clearwater Rivers. The Salmon and Clearwater Rivers flow westward through central Idaho. The

Salmon River flows into the Snake River south of Lewiston, Idaho, while the Clearwater River flows into the Snake River at Lewiston.

Researchers usually focus on the supposed ancestral Columbia River. They line up the quartzite gravels in the southwest Columbia Basin east of the Cascades from Sentinel Gap southwest across the Horse Heaven Hills at Satus Pass to Hood River, Oregon (figure 11). They then suggest that a very wide paleovalley occupied by the ancestral Columbia River flowed along this route and that Satus Pass and other wind gaps along some of the other ridges represent old river channels.^{43–46} There are several wind gaps through the Rattlesnake Hills that have been attributed to the ancestral Columbia River. However, quartzites are *not* found in these wind gaps as one would expect, but commonly on the highest terrain of the ridges with *no* indication of an ancient channel, as for instance on the Horse Heaven Hills.⁴⁷ There are quartzites on the ridge east of the low spot at Satus Pass, considered a wind gap, but none in Satus Pass itself.⁴⁸ Such missing quartzites that are expected in wind gaps led Aaron Waters to reject this so-called ancestral path of the Columbia River.⁴⁹

Waters also pointed out that quartzite-bearing conglomerates are widespread as a thin, discontinuous veneer throughout south-central Washington and northern Oregon. The quartzite gravels in Washington are found over too broad an area to ascribe them to some ancestral river. There are not only quartzite gravels found widely over the south-central Washington Columbia Basin, including the tops of basalt ridges, but they are also found along the ridges of the Columbia Gorge. Quartzite gravels are also found in many other areas of eastern Washington, such as in southeast Washington under the Palouse silt and near

Cheney, southwest of Spokane. There does not appear to be any evidence to suggest the remnants of a great ‘ancestral river’ bed. Moreover, the ‘ancestral river’ hypothesis does not account for the quartzite gravels claimed to be on high ridges on the eastern side of the Cascades⁵⁰ or those in the Puget Sound area. The distribution suggests something more like a sheet flow of quartzite gravel followed by the later formation of anticlines in the basalt and erosion into remnants. Besides, it is doubtful there is enough stream power to transport quartzite from the Rocky Mountains to the Pacific Ocean, especially considering the sinuous courses of any supposed paleoriver.³

We now turn our attention to northern and central Oregon. Allen⁵¹ and others noted at least eight locations of well-rounded quartzite boulders with abundant percussion marks on *top* of the Wallowa Mountains, northeast Oregon, from 1,402 to 2,658 m ASL. The boulders are up to 1 m in diameter southeast of Lookout Mountain, and Jim White Ridge is covered by a 55-m thick layer of quartzite boulders up to 60 cm in diameter, covering an area of about 3 km². Placer gold exists in the quartzite gravels and an active mine once existed on Jim White Ridge.

Allen first discovered these gravels in 1938 but published little on the subject until 1991 because they puzzled him. He suggested that the gravels with gold in the matrix were not laid down by some lazy meandering stream flowing from the east or southeast, but by a northwest flowing ‘*torrential paleoriver*’ in a broad valley with a braided channel (figure 12).⁵¹ The question is what paleoriver? The river apparently is *not* the ancestral Columbia or Snake Rivers.⁵² (We suppose that the uniformitarian geologists can claim the quartzite gravels were spread by the ancestral Clearwater or Salmon Rivers, when the Snake River supposedly flowed

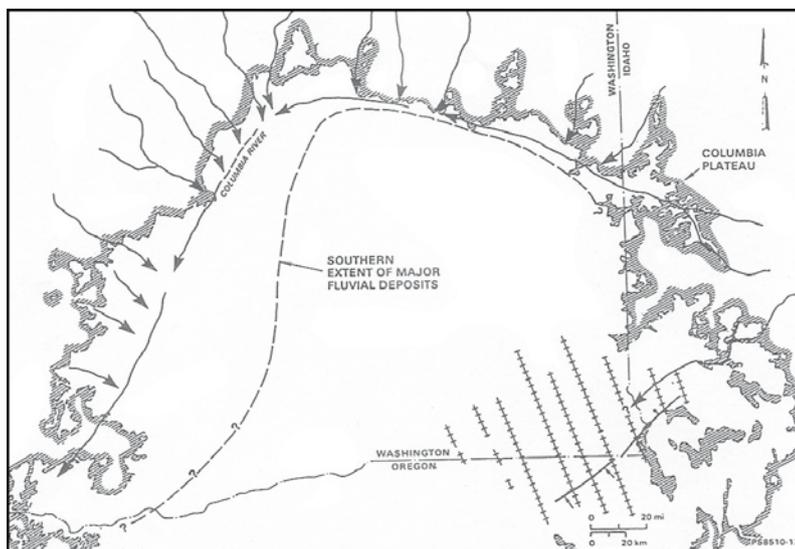


Figure 11. The supposed course of the ‘ancestral’ Columbia River along the edge of the Columbia River basalts. The path through the southwest Columbia River Basalt Group based on the deposited quartzite gravels from Sentinel Gap southwest to Hood River, Oregon. (From Fecht *et al.*,⁴² p. 226).

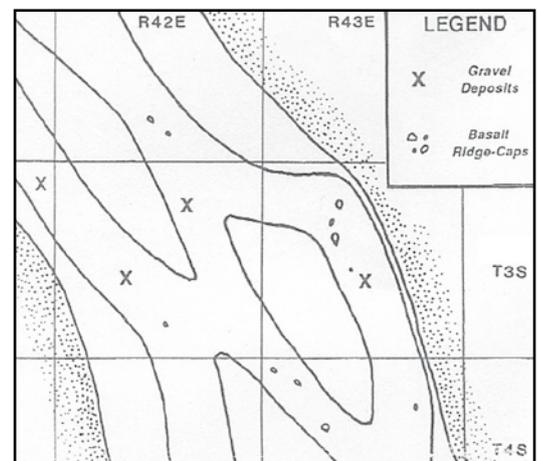


Figure 12. Schematic of a wide northwesterly flowing, braided stream through the northwest part of the Wallowa Mountains in early Miocene time according to Allen.⁵⁹ Stipples indicate walls of valley and x marks locations where quartzite gravels are found high up in the Wallowa Mountains. Periodic torrential flow spreads the quartzites farther and farther northwest.

through southeast Oregon and before the Snake cut the deepest water gap in North America at Hells Canyon. There is little if any evidence for the existence of a paleo-Snake River through Oregon.^{53,54} Allen admits that his deduction is ‘an outrageous hypothesis’, but it is the best he can do, and he published the information in 1991 to stimulate further research.

The nearest source of quartzite outcrops is in central Idaho, about 200 km to the east, but across Hells Canyon. Our research has revealed other quartzite gravel locations several km west of those that Allen cites. There are even more in northeast Oregon that we have not seen but are briefly mentioned in the literature.^{55,56} Their presence in more locations spells an even more widespread distribution than Allen mentions and makes his broad paleovalley hypothesis even more outrageous. We believe Allen was on the right track in deducing a paleotorrent in a wide channel,⁵¹ but the quartzites are too extensive to provide a hypothesis for just river transport in the area of the Wallowa Mountains before they rose.

The quartzite locations in central Oregon on top of Gold Hill and east of Paulina are especially difficult to place within the confines of some ‘ancestral river’. They are well southwest of the supposed paths of the ancestral Clearwater or Salmon Rivers. There is no evidence in the geology of central or eastern Oregon to suggest an ancient river.

Conclusion

The uniformitarians themselves recognise the problems associated with the sheer scale of the distribution of quartzites either side of the Rocky Mountains.⁵¹ The main uniformitarian hypothesis seems to be fluvial transport, but this hypothesis falls far short of accounting for the long distance of transport, the iron staining, and percussion marks.³ Other previous explanations offered for the eastward transport of quartzites, such as transport by tectonic uplift in eastern Idaho and transport on thrust sheets that have moved east for a long distance have been discarded by uniformitarians. None others were offered for long distance transport east of the source.

The major problem with the uniformitarian explanations of the quartzites is that they do not have access to geological processes that have sufficient power to explain the extent and thickness of the quartzite deposits and the massive scale water volume and velocity that must inevitably go with it. The failure of these points amounts to a failure of the interpretive framework, and the need for one that has sufficient power to explain the data.

We believe, however, that the quartzite gravels represent strong evidence for the Recessive Stage of the Flood with its two phases of sheet flow followed by channelized flow.⁵⁷ In the final paper of this series,⁵⁸ we explore a diluvial interpretation of these quartzites within the context of the Recessive Stage of the Flood, and show how it explains the evidence in an elegant way.

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